

A new biostratigraphically significant calcareous nannofossil species in the Early Pliocene of the Mediterranean

Carlos Lancis¹ and José-Abel Flores²

¹Departamento de Ciencias de la Tierra, Facultad de Ciencias, Universidad de Alicante, Spain

²Departamento de Geología, Facultad de Ciencias, Universidad de Salamanca, Spain

email: flores@usal.es

ABSTRACT: Here we describe *Reticulofenestra cisnerosii* nov. sp. a rounded species of “Reticulofenestrid” with a well defined stratigraphic distribution in the Mediterranean Neogene. The first occurrence of this new taxa is placed in the middle part of subchron Thvera (C3n4n); in the subzone CN10b (Okada and Bukry 1980); and the last occurrence is observed into the subchron C3n1r, middle part of subzone CN 11a (Okada and Bukry 1980), just before the regular record of *Discoaster asymmetricus*. In the region studied, this species is linked with warm, moderately saline and relatively fertile waters. This species has also been observed in the equatorial Pacific.

INTRODUCTION

Reticulofenestra is a common genus in the geological record through the Paleogene and Neogene (including the Quaternary). Several events based on different species of *Reticulofenestra* have been used as biostratigraphic markers (Backman 1980; Pujos 1987; Driever 1988; Gallagher 1989; Young 1990; Rio et al 1990; Raffi and Flores 1995; Kameo and Takayama 1999; Okada, 2000; Marino and Flores 2002, among others). Lancis (1998) carried out a qualitative and quantitative study of Neogene materials from the Betic region (figure 1) mainly based on Reticulofenestrids, with a focus on biostratigraphic and paleoecological reconstructions. Lancis (1998) used both standard and regional biozonations (Martini 1971; Okada and Bukry 1980; Rio et al. 1990; Raffi and Flores 1995; Fornaciari et al. 1996) for calibration of several newly defined events. One of the most characteristic events in the lower Pliocene is the record of rounded forms of *Reticulofenestra* (from 4 to 8µm) which occur just after the first occurrence of *Ceratolithus acutus* (Gartner and Bukry 1974), in the Messinian-Pliocene interval. These forms are clearly different from *Pseudoemiliania lacunosa* (Kamptner 1963; Gartner 1969) and *Reticulofenestra pseudoumbilicus* (Gartner 1967; Gartner 1969), both of which are abundant in this interval. Morphological and stratigraphic criteria are used to distinguish these specimens from others with similar characteristics.

TAXONOMY

CLASS Prymnesiophyceae Hibberd 1976

FAMILY Noelaerhabdaceae Jerkovic 1970

GENUS *Reticulofenestra* Hay, Mohler and Wade 1966 emend.

Gallagher 1989

Reticulofenestra cisnerosii Lancis and Flores n. sp.

Plate 1, figures 1-10

Etymology: “In memoriam” Professor Daniel Jiménez de Cisneros y Hervás teacher in *Instituto de Segunda Enseñanza* of Alicante between 1904 and 1933 and pioneer palaeontologist in Alicante (Spain).

Description: Circular to subcircular placoliths with a well developed collar with a distal part of proximal shield elevated, defining a prominent collar with high brightness when observed under crossed nichols. The central opening constitutes at least one third of the total diameter. The distal shield is similar in thickness at the collar but less birefringent. Periphery is slightly jagged.

Size between 4 and 8µm.

Occurrence: *Reticulofenestra cisnerosii* have their first occurrence in the early Pliocene, after the first occurrence of *Ceratolithus acutus*, in the Thvera chron (C3n4n- middle part of Subzone CN10b); and their last occurrence in the middle part of Subzone CN11a, just before the first record (scattered) of *Discoaster asymmetricus* Gartner 1969 in chron C3n4n. Based on Berggren et al. (1995), this represents an interval between 5.23 and 4.48 Ma.

Preliminary studies carried out in the equatorial Pacific (Ocean Drilling Project Site 202) show a similar distribution.

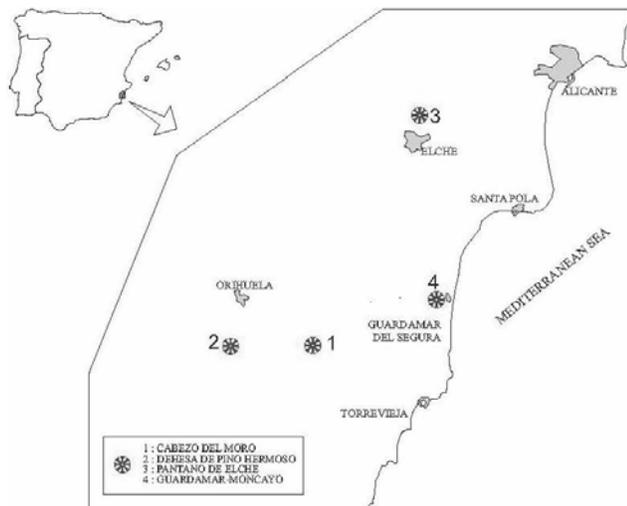
Type locality: Dehesa de Pino Hermoso. (Alicante) Spain.

Section DPH. Holotype: sample DPH20 (fig. 1a,b)

Paratypes: Samples DPH 14 (fig. 2a,b; 3a,b), DPH 15 (fig. 6), PTEL 4 (fig. 4a,b), PTEL 6 (fig 7-10), PTEL 9 (fig. 5a,b).

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Remarks: A narrow distal shield with a prominent collar as well as circular-subcircular form allow us to differentiate *R. cisnerosii* from *R. pseudoumbilicus*. *Reticulofenestra rotaria* (Theodoridis 1984=*Reticulofenestra pseudoumbilicus* var. *rotaria* sensu Young 1990) is similar in morphology, but this species is characteristic from the Messinian in the Betic sections. *Reticulofenestra asanoi* (Sato and Takayama 1992) is also similar to *R. cisnerosii* but their record is exclusively Pleistocene (Sato and Takayama 1992; Wei 1993). Kameo and Takayama (1999) and Kameo and Bralower (2000) observed in Deep Sea



TEXT-FIGURE 1
Geographic situation of the studied sections.

Drilling Program Sites 606, 607 and 610 (North Atlantic) and ODP Sites 998, 999 and 1000 (Caribbean Sea), respectively, a similar morphology named *R. minutula* var. A (circular form) larger than 5 mm, but they occur in the lower part of zone CN12a. *R. cisnerosii* is easy to distinguish from *Reticulofenestra pacifica* Nishida 1979 due to their elliptical form.

Paleoecological remarks

Based in the nannofossil assemblage, Lancis (1998) interpreted *R. cisnerosii* (morphotypes now named in this way) as being more abundant in warm, moderately saline and relatively fertile waters, due to low proportions of *R. pseudoumbilicus* and high proportions of small and medium *Reticulofenestra* (*R. haqii/minutula* and *R. minuta*) and *Helicosphaera*, as well as and regular presence of sphenoliths, asteroliths and *Calcidiscus* spp.

Comments

As noted, the presence of circular species is recurrent during the late Neogene. In several cases, such as with *R. rotaria*, these morphologies are frequent to abundant under conditions of ecological stress in semi-closed basins (Flores et al., 2005), although they are also present in open ocean (Raffi and Flores 1995) in minor proportions. This species is abundant at the base of the Pliocene in the Mediterranean region immediately after the Messinian Salinity crisis, returning to normal oceanic conditions but still under certain restriction due their semi-closed basin characteristic. At this time an intense seasonality under relatively warm conditions were the dominant environmental conditions (Lancis 1998). Other taxa, such as *R. minutula* var. A or *R. asanoi*, developed this rounded morphology synchronously in the ocean.

This intriguing evolutionary behaviour raises several questions: Do these forms respond to a common strategy? Can they be considered ecophenotypic modifications? How can this study be approached?

For us, the repetition of this circular morphology represents a case of iterative evolution, from the common branch of *R.*

pseudoumbilicus. This process occurs at different intervals: Messinian, lower Pliocene, late Pliocene and middle Pleistocene, responding to equivalent selective pressure. Stressing conditions where intense seasonality could play a crucial role favours the development of circular forms. These changes seem to occur in relatively simple ecophenotypic features such as the coccolith morphology. These kinds of characteristics are probably controlled by a reduced number of genes with a common regulation. The appearance of weak modifications in the coccolith morphology or in the amount of calcite produced in the collar only would imply punctual mutations with high probability of success. The possibility of selection of these changes is linked to favouring environmental conditions. The repetition of these conditions towards the Neogene seems to be the reason of this iteration.

We note that, however, that specimens with intermediate morphology between *R. cisnerosii* and *P. lacunosa* (occasional presence of few small windows in the distal shield) have been observed. These characteristics, strictly morphological, raise the possibility that some forms of *P. lacunosa* evolved from *R. cisnerosii*. In any case, the morphological variations in shape and size observed in this species (e.g., Young 1990; Driever 1988) suggest a polyphyletic origin. Additional studies are necessary to clarify these ideas.

We are conscious that this is the beginning of a challenge to decipher this peculiar variability, but at the same time we consider that identification and calibration of these forms in the sedimentary sequence constitutes a useful biostratigraphic tool.

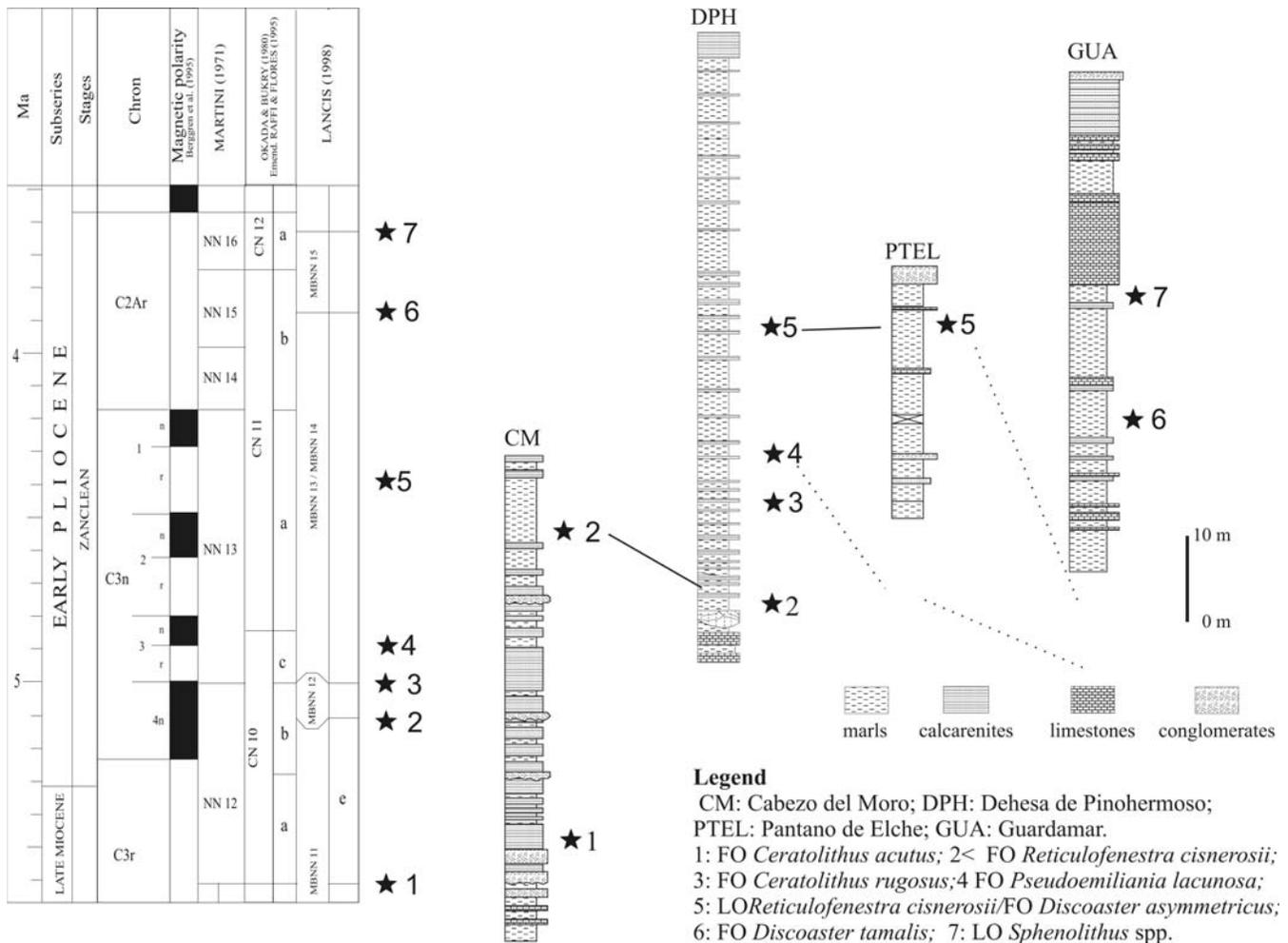
ACKNOWLEDGMENTS

We would like to thank Dr. Mario Cachao (University of Lisbon, Portugal), Dr. Lluïsa Cros (Institute for Marine Research-CSIC, Barcelona, Spain) and Dr. Gabriel Filippelli (IUPUI, USA) for their comments and discussion during the preparation of this manuscript.

This study was supported by Instituto de Cultura Juan Gil-Albert (Alicante, Spain) and research grant CLI-1002-CO2-02 (Ministerio de Educación y Ciencia)

REFERENCES

- BACKMAN, J., 1980. Miocene-Pliocene nannofossils and sedimentation rates in the Hatton-Rockall basin, N.E Atlantic ocean. *Act. Univ. Stockholm Contributions to Geology*, 36(1): 1-93.
- BERGGREN, W.A., KENT, D.V., SWISHER, C.C. and AUBRY, M.P., 1995. A revised Cenozoic geochronology and chronostratigraphy. In: Berggren, W. A., Kent, D. V., Aubry, M.-P. and Harebell, I., *Geochronology time scales and global stratigraphic correlation*, 129-212. Tulsa: SEPM Special Publication 54.
- DRIEVER, B.W.H., 1988. Calcareous nannofossil biostratigraphy and Paleoenvironmental interpretation of the Mediterranean Pliocene. *Utrecht Micropaleontological Bulletin*, 36: 1-245 p.
- FLORES, J. A., SIERRA, F.J., FILIPPELLI, G., BÁRCENA, and M.A. PÉREZ-FOLGADO, M. 2005. Surface water dynamics and biogeochemical cycling during deposition of cyclic late Messinian sapropel sequences in the western Mediterranean. *Marine Micropaleontology*, 54 (1-2), 50-79.
- FORNACIARI, E., DI STEFANO, A., RIO, D., and NEGRI, A. Middle Miocene quantitative calcareous nannofossil biostratigraphy in the Mediterranean region. *Micropaleontology* 42:37-63



TEXT-FIGURE 2

Stratigraphic correlation panel. Lithology and situation of calcareous nanofossil biostratigraphy events: FO and LO of *Reticulofenestra cisnerosii* nov.sp. and their relationship with the geomagnetic and calcareous nanofossil biostratigraphy standard scales. FO- first occurrence; LO- last occurrence.

GALLAGHER, L., 1989 *Reticulofenestra*: A critical review of taxonomy, structure and evolution. In: Crux, J.A. and Van Heck, S.E., Eds., *Nanofossils and their applications*, 41-75. British Micropaleontological Society Series. Chichester: Ellis Horwood Ltd.

HIBBERD, D.J., 1976. The ultrastructure and taxonomy of the Chrysophyceae and Prymnesiophyceae (Haptophyceae): a survey with some new observations on the ultrastructure of the Chrysophyceae. *Botanical Journal of the Linnean Society*, 72: 55-80.

JERCOVIC, L., 1970. *Noelaerhabdus* nov. gen. Type d'une nouvelle famille de coccolithophoridés fossiles. Noelaerhabdaceae du Miocène supérieur de Yougoslavie. *Comptes Rendus, Hebdomad des Seances Academie Science*, 270: 468-470.

KAMEO, K. and TAKAYAMA, T., 1999. Biostratigraphic significance of sequential size variations of the calcareous nanofossil genus *Reticulofenestra* in the Upper Pliocene of the North Atlantic. *Marine Micropaleontology* 37:41-52.

KAMEO, K. and BRALOWER, J., 2000. Neogene calcareous nanofossil biostratigraphy of sites 998, 999, and 1000, Caribbean Sea. *Proceedings of the Ocean Drilling Program, Scientific Results*, 165: 3-17. College Station, TX: Ocean Drilling Program.

LANCIS, C., 1998. "El Nanoplancton Calcáreo de las Cuencas Neógenas Orientales de la Cordillera Bética." Doctoral Thesis, Universite de Alicante (España). 841p.

MARINO, M. and FLORES, J.A., 2002. Miocene to Pliocene calcareous nanofossil biostratigraphy at ODP Leg 177 Sites 1088 and 1090. *Marine Micropaleontology* 45:291-307

MARTINI, 1971. Standard Tertiary and Quaternary calcareous nanoplankton zonation. In: Farinacci, A., Ed., *Proceedings, Second International Conference on Planktonic Microfossils, Roma*, Volume 2, 739-785. Rome: Edition Tecnoscienza.

NISHIDA, S., 1979. Atlas of Pacific Nanoplanktons. *News of Osaka Micropaleontologists*, Special Paper, No. 3, 31p.

- OKADA, H., 2000. Neogene and Quaternary calcareous nannofossils from the Blake Ridge, Sites 994, 995 and 997. *Proceedings of the Ocean Drilling Program, Scientific Results*, 164: 331-341. College Station, TX: Ocean Drilling Program.
- OKADA, H. and BUKRY, D., 1980. Supplementary Modification and introduction of code numbers to the Low-Latitude coccolith biostratigraphic zonation (Burkry 1973-1975). *Marine Micropaleontology*, 5: 321-325.
- PUJOS, A., 1987. Late Eocene to Pleistocene medium-sized and small-sized "Reticulofenestrids." *Abhandlungen der Geologischen Bundesanstalt*, 39: 239-277.
- RAFFI, I. and FLORES, J.A., 1995. Pleistocene through Miocene Calcareous nannofossils from eastern equatorial Pacific Ocean (Leg 138). *Proceedings of the Ocean Drilling Program, Scientific Results*, 138: 233-285. College Station, TX: Ocean Drilling Program.
- RIO, D., FORNACIARI, E., and RAFFI, I., 1990. Late Oligocene through early Pleistocene calcareous nannofossils from western equatorial Indian Ocean (Leg 115). In: Duncan, R.A., Backman, J., Peterson, L.C., et al., *Proceedings of the Ocean Drilling Program, Scientific Results*, 115:175-235. College Station, TX: Ocean Drilling Program.
- SATO, T. and TAKAYAMA, T., 1992. A stratigraphically significant new species of the calcareous nannofossil *Reticulofenestra asanoi*. In: Ishizaki, K. and Sato, T., Eds., *Centenary of Japanese Micropaleontology*, 457-460. Tokyo: Terra Scientific Publishing Company.
- THEODORIDIS, S. A., 1984. Calcareous nannofossils biozonation of the Miocene and revision of the Helicoliths and Discoasters. *Utrecht Micropaleontological Bulletin*, 31: 1-271.
- WEI, W., 1993. Calibration of upper Pliocene-lower Pleistocene nannofossil events with oxygen isotope stratigraphy. *Paleoceanography* 8, 85-99.
- YOUNG, J.R., 1990. Size variation of Neogene Reticulofenestrid coccoliths from Indian Ocean DSDP Cores. *Journal of Micropaleontology*, 9:71-85.
- Manuscript received August 21, 2006
Manuscript accepted September 11, 2006

PLATE 1

Reticulofenestra cisnerosii. Magnification $\times 1500$.
1-5, Petrographic microscope: a, parallel nichols and b crossed nichols.

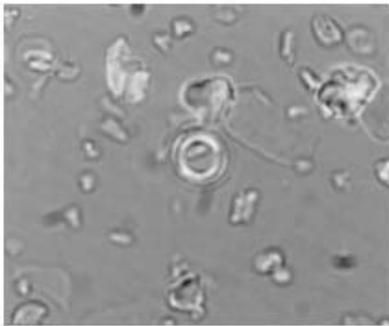
1a,b Holotype; sample DPH 20.

4a,b Paratype; sample PTEL 6.

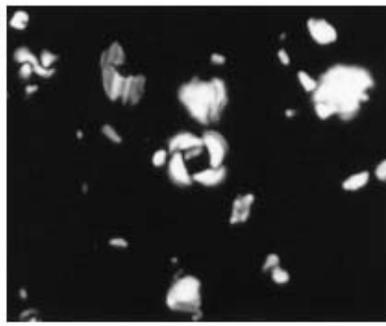
2a,b Paratype; sample DPH 14.

5a,b Paratype; samples PTEL 9.

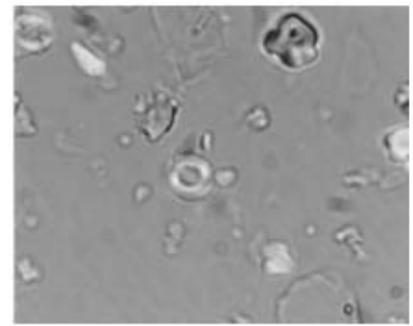
6-10 — Scanning electron microscope. Paratype; sample PTEL 6.



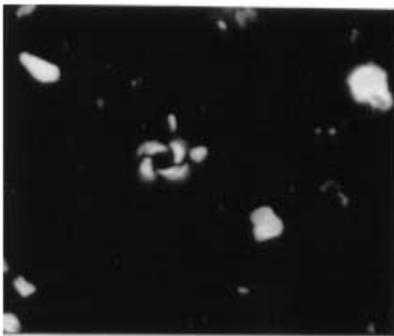
1a



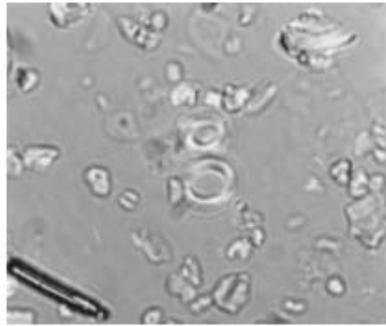
1b



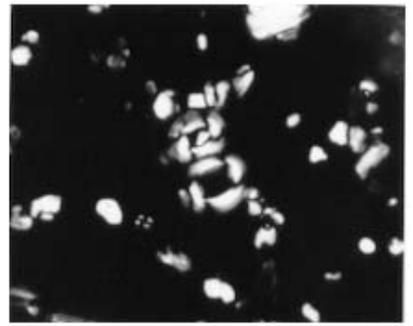
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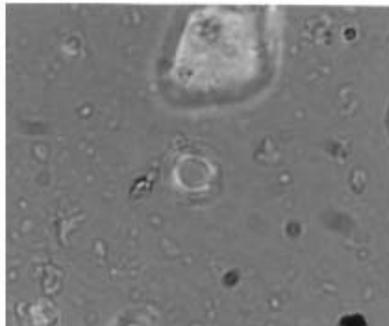
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3a



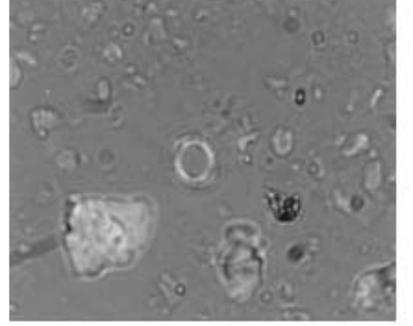
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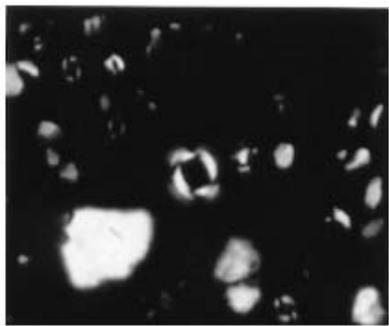
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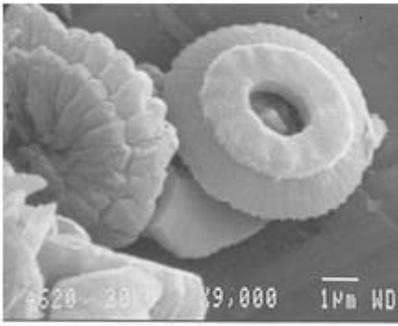
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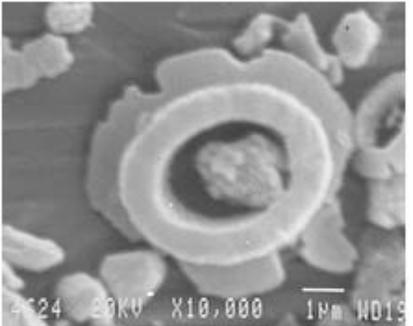
5a



5b



6



7